

# ラジカル素反応過程と 大気と燃焼の科学

東京大学 大学院工学系研究科 化学システム工学専攻

三好 明

「プラズマ科学における分光計測の高度化と原子分子過程研究の新展開」  
「原子分子データ応用フォーラムセミナー」合同研究会 (2016年12月21日 核融合研究所/土岐)

## 私 (三好 明)

専門：化学反応論・燃烧化学・大気化学

1990 工学博士 (東京大学・反応化学専攻)

1990-環境庁 国立環境研究所 大気圏環境部 研究員

1992-東京大学 工学部 反応化学科 助手

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准教授

詳細：<http://www.frad.t.u-tokyo.ac.jp/>

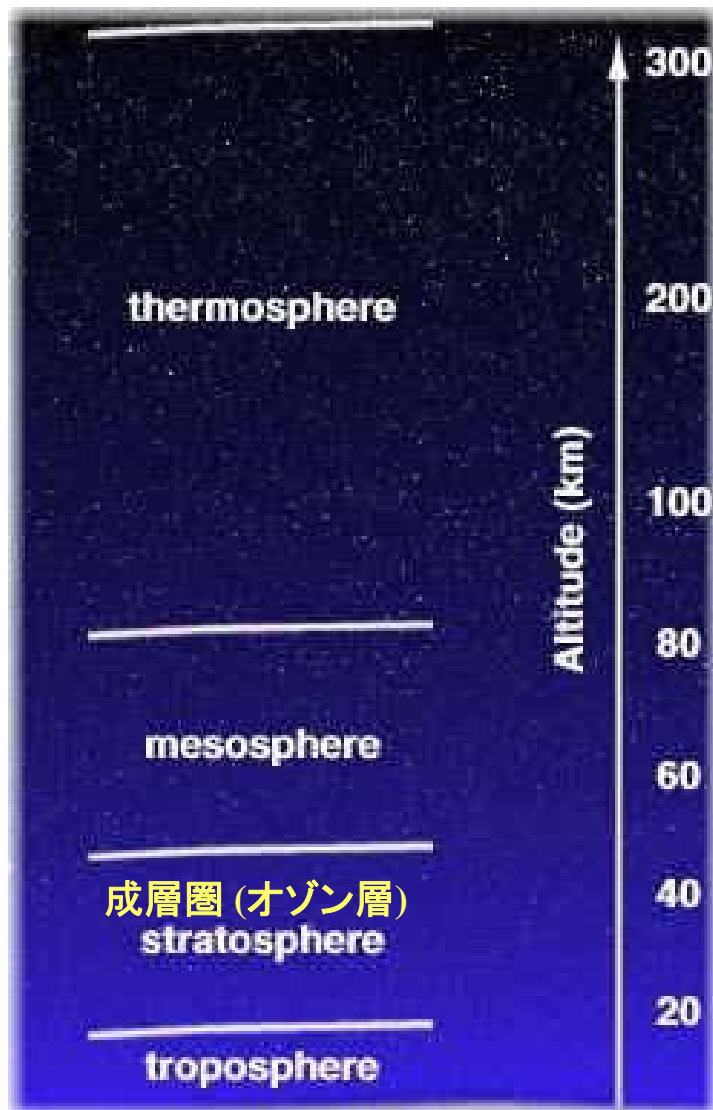
# オゾン層破壊

## Ozone Depletion

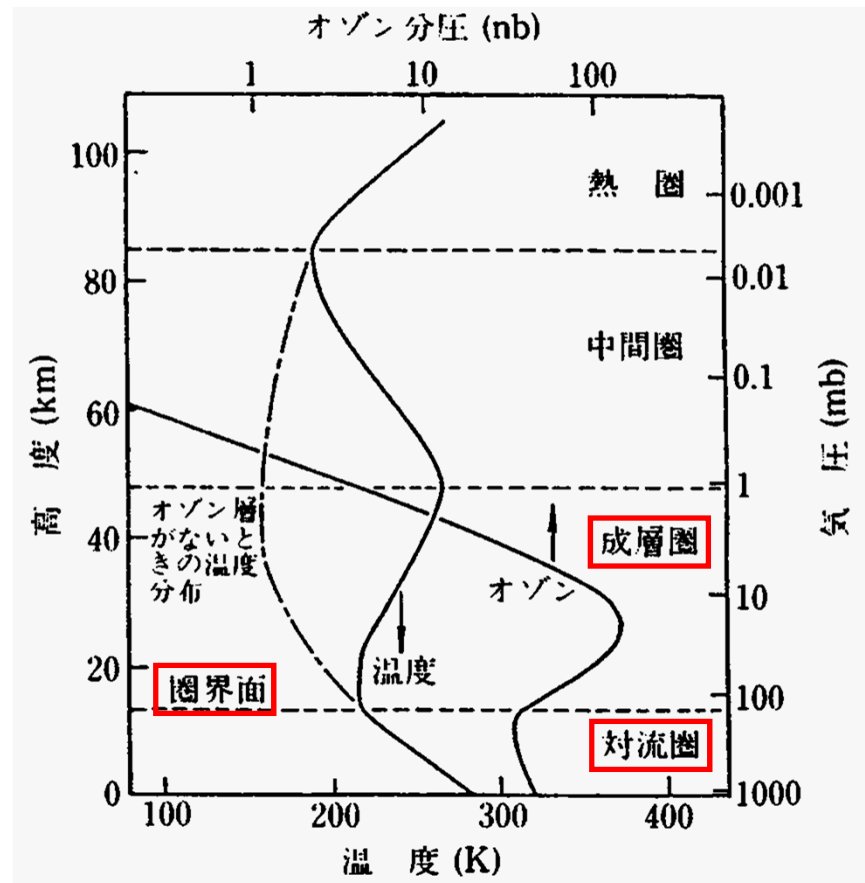
— 地球規模環境問題への  
対策の優等生



# 地球大気の構造

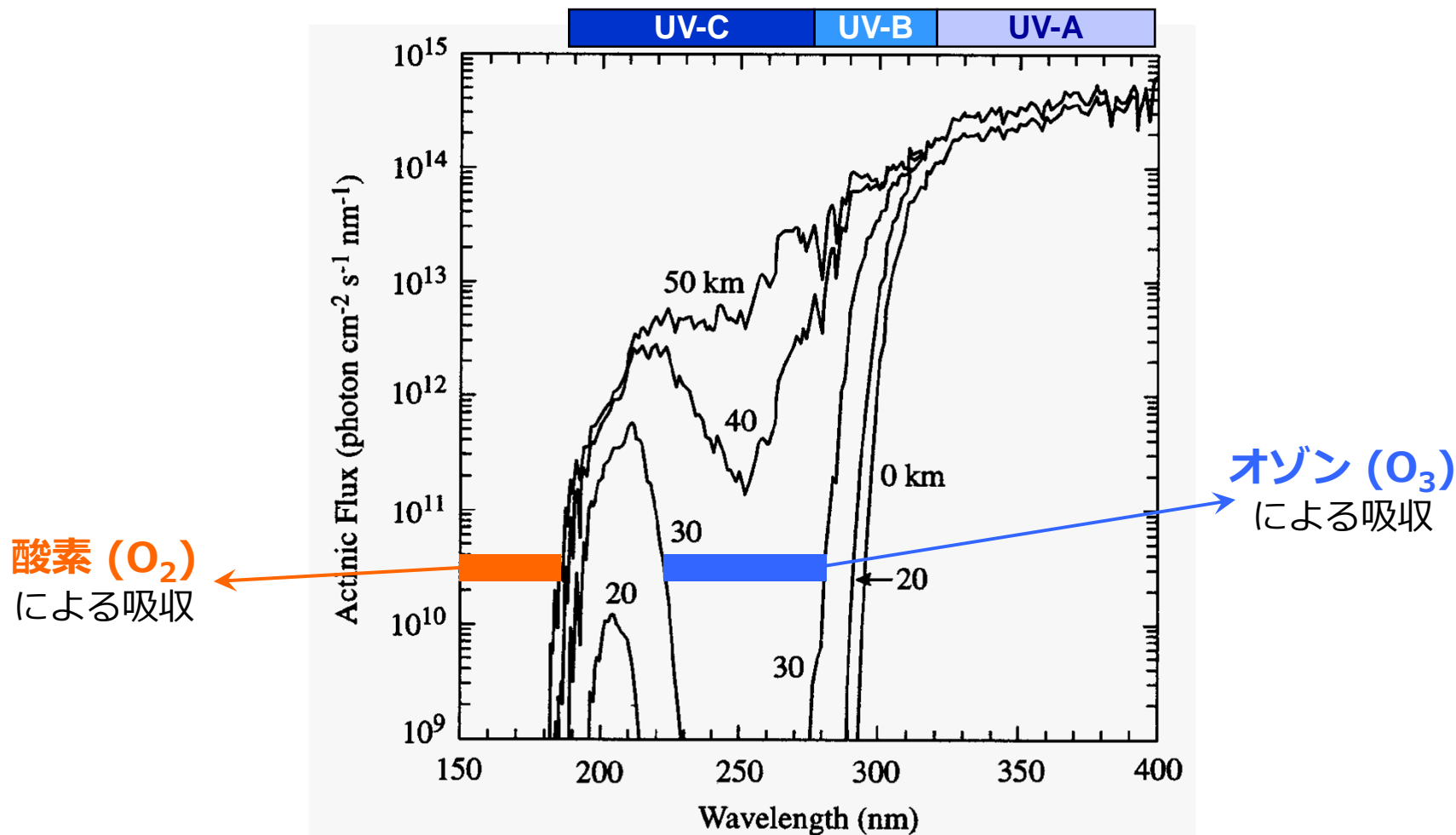


大気の >99% は成層圏・対流圏に存在



"フロン 地球を蝕む物質," 富永 健, 卷出 義紘,  
F. S. Rowland, 東大出版会, 東京 (1990).

# 太陽光の波長分布



**Fig. 10-2** Solar actinic flux at different altitudes, for typical atmospheric conditions and a 30° solar zenith angle. From DeMore, W. B., et al. *Chemical Kinetics and Photochemical Data for Use in Stratospheric Modeling*. JPL Publication 97-4. Pasadena, Calif.: Jet Propulsion Lab, 1997.

# CFC (クロロフルオロカーボン)

## 1961 Lovelock – ECD 開発

ハロゲン・硫黄を含む化合物の高感度検出器  
[Anal. Chem., 33, 162 (1961)]

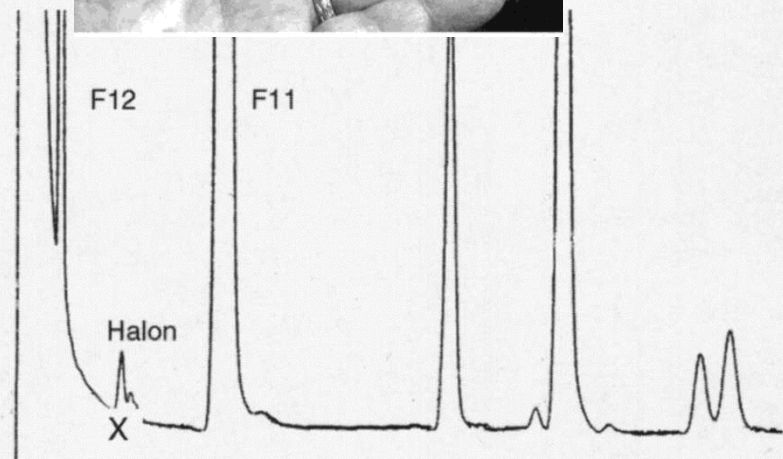


Figure 1. A chromatogram to illustrate the sensitivity of the ECD. The peak 'X' represents 1.1 parts per trillion of Halon in the air

J. Lovelock, "Homage to Gaia," Oxford Univ. Press, Oxford (2000).

## 1971 Lovelock – 大気中 CFC-11

天然には存在しない  $\text{CCl}_3\text{F}$  を大気中で測定  
[Nature, 230, 379 (1971)]

NATURE VOL. 230 APRIL 9 1971

### Atmospheric Fluorine Compounds as Indicators of Air Movements

GASEOUS fluorine compounds are supposed not to occur naturally in the atmosphere. Volatile fluorine compounds would not be expected to result from chemical equilibria between fluorine compounds on the surface of the Earth, and it is improbable that biological systems contribute significant quantities of organic fluorine compounds.

Table 1 Observations at Adrigole, Co. Cork, Ireland (51° 40' N, 09° 45' W)

Wind heading	Concentration by volume		Turbidity
	$\text{CCl}_3\text{F}$	$\text{SF}_6$	
45°-135°	$1.0 \times 10^{-11}$ (4)	$2.9 \times 10^{-14}$ (3)	0.03 (7)
225°-315°	$1.9 \times 10^{-10}$ (3)	$1.2 \times 10^{-13}$ (3)	0.19 (2)

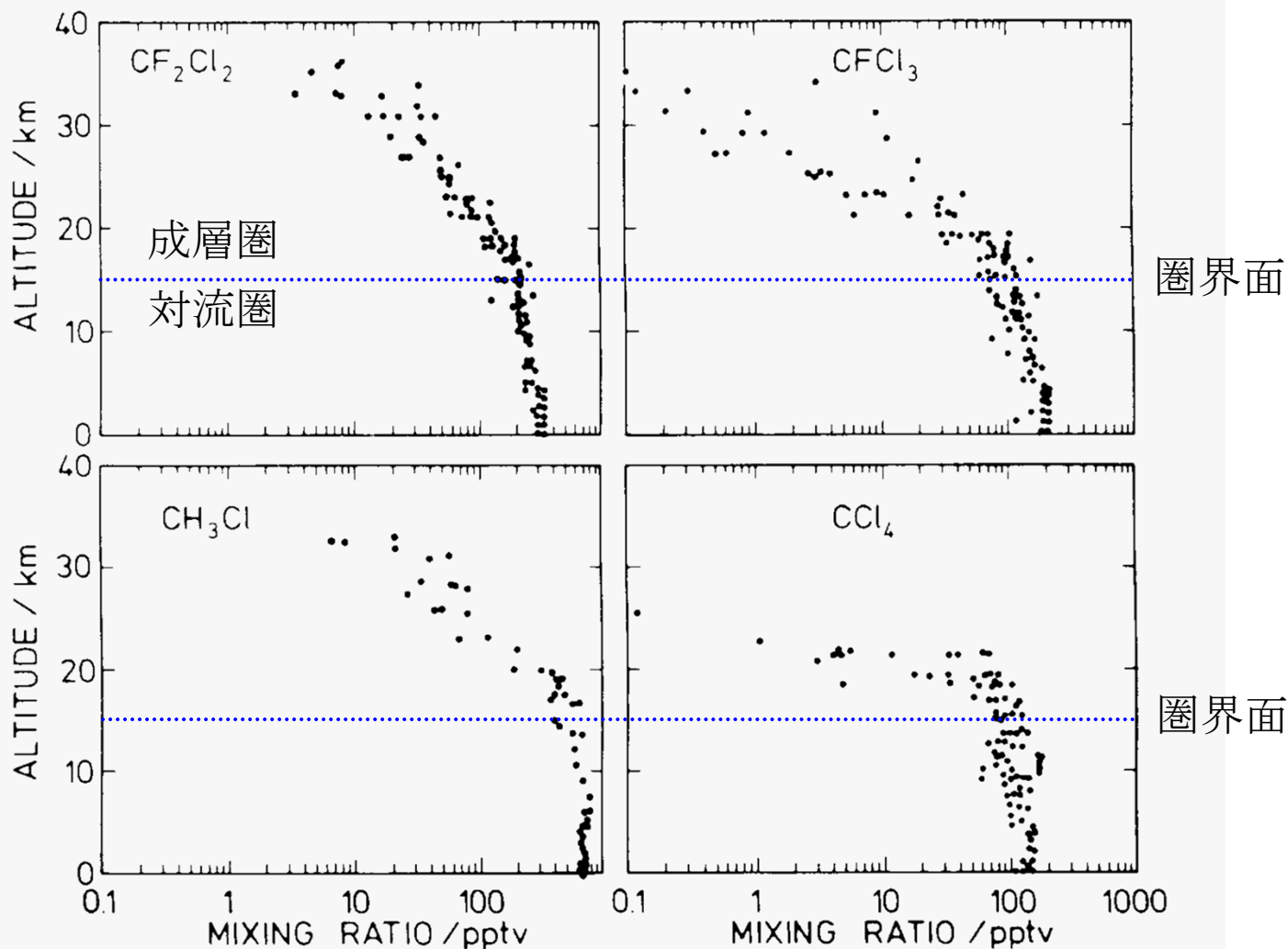
The number of observations is shown in parentheses.

J. E. LOVELOCK

Department of Applied Physical Science,  
University of Reading,  
Reading RG6 2AL

# 大気中の CFC (クロロフルオロカーボン)

## 大気中 CFC の鉛直分布



圏界面 (~15 km) で急激な減少 → 成層圏で分解

"Chemistry of the Natural Atmosphere," P. Warneck, Academic Press, San Diego (1988).

## 1974 Molina &amp; Rowland – オゾン層破壊予言

[Nature, 249, 810 (1974)] (1995ノーベル化学賞)

810

Nature Vol. 249 June 28 1974

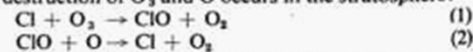
## Stratospheric sink for chlorofluoromethanes : chlorine atom-catalysed destruction of ozone

Mario J. Molina & F. S. Rowland

Department of Chemistry, University of California, Irvine, California 92664

*Chlorofluoromethanes are being added to the environment in steadily increasing amounts. These compounds are chemically inert and may remain in the atmosphere for 40–150 years, and concentrations can be expected to reach 10 to 30 times present levels. Photodissociation of the chlorofluoromethanes in the stratosphere produces significant amounts of chlorine atoms, and leads to the destruction of atmospheric ozone.*

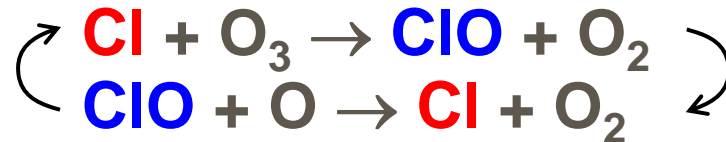
photolytic dissociation to  $\text{CFCl}_2 + \text{Cl}$  and to  $\text{CF}_2\text{Cl} + \text{Cl}$ , respectively, at altitudes of 20–40 km. Each of the reactions creates two odd-electron species—one Cl atom and one free radical. The dissociated chlorofluoromethanes can be traced to their ultimate sinks. An extensive catalytic chain reaction leading to the net destruction of  $\text{O}_3$  and O occurs in the stratosphere:



This has important chemical consequences. Under most conditions in the Earth's atmosphere the reactions are slower because of the lower concentrations of the reactants.



連鎖反応

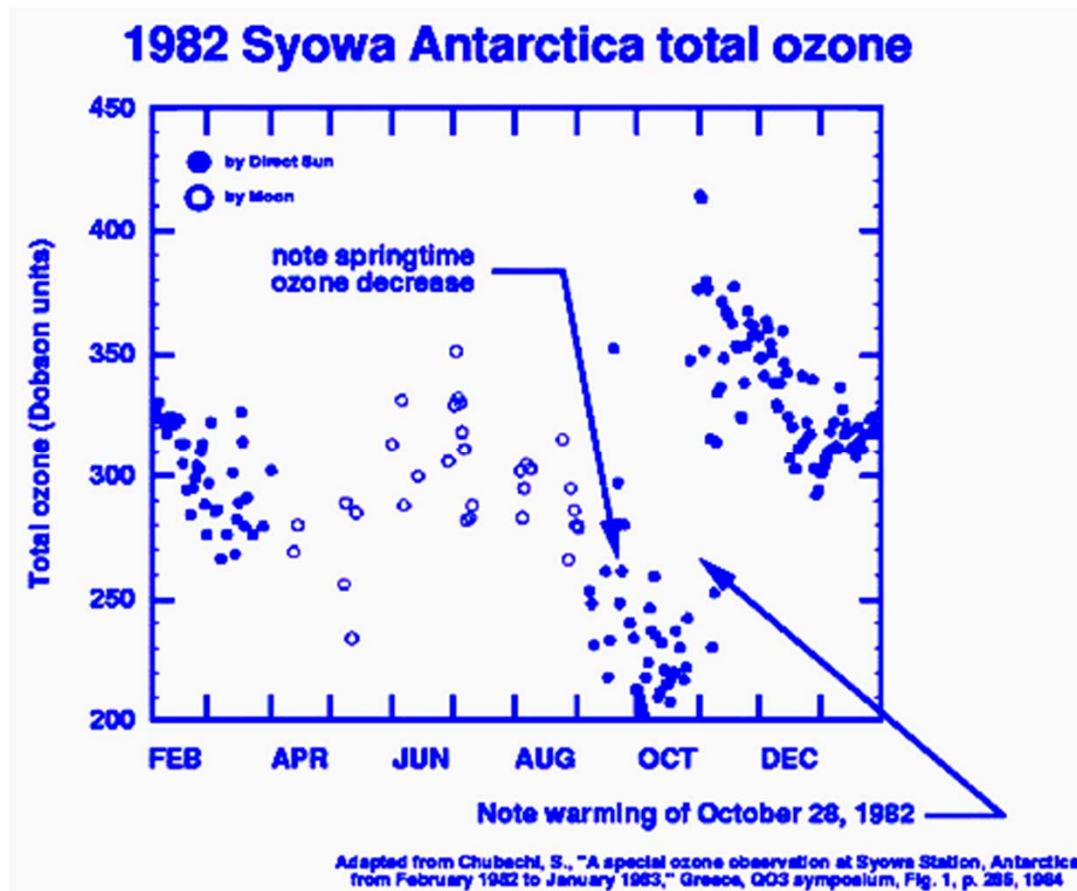


1 つの Cl 原子が  $\sim 10^4$  個  
のオゾン ( $\text{O}_3$ ) を破壊





## 1984 Chubachi – 南極オゾン異常減少 (1982)



cf.) 100 DU  
= 1mm O<sub>3</sub> STP

S. Chubachi, "A special ozone observation at Syowa Station, Antarctica from February 1982 to January 1983," Greece, QO3 symposium, Fig. 1, p. 286, 1984

# 南極オゾンホール

≠ オゾン層破壊

1985 Farman et al. – 南極オゾンホール  
[Nature, 315, 207 (1985)]

NATURE VOL. 315 16 MAY 1985 207

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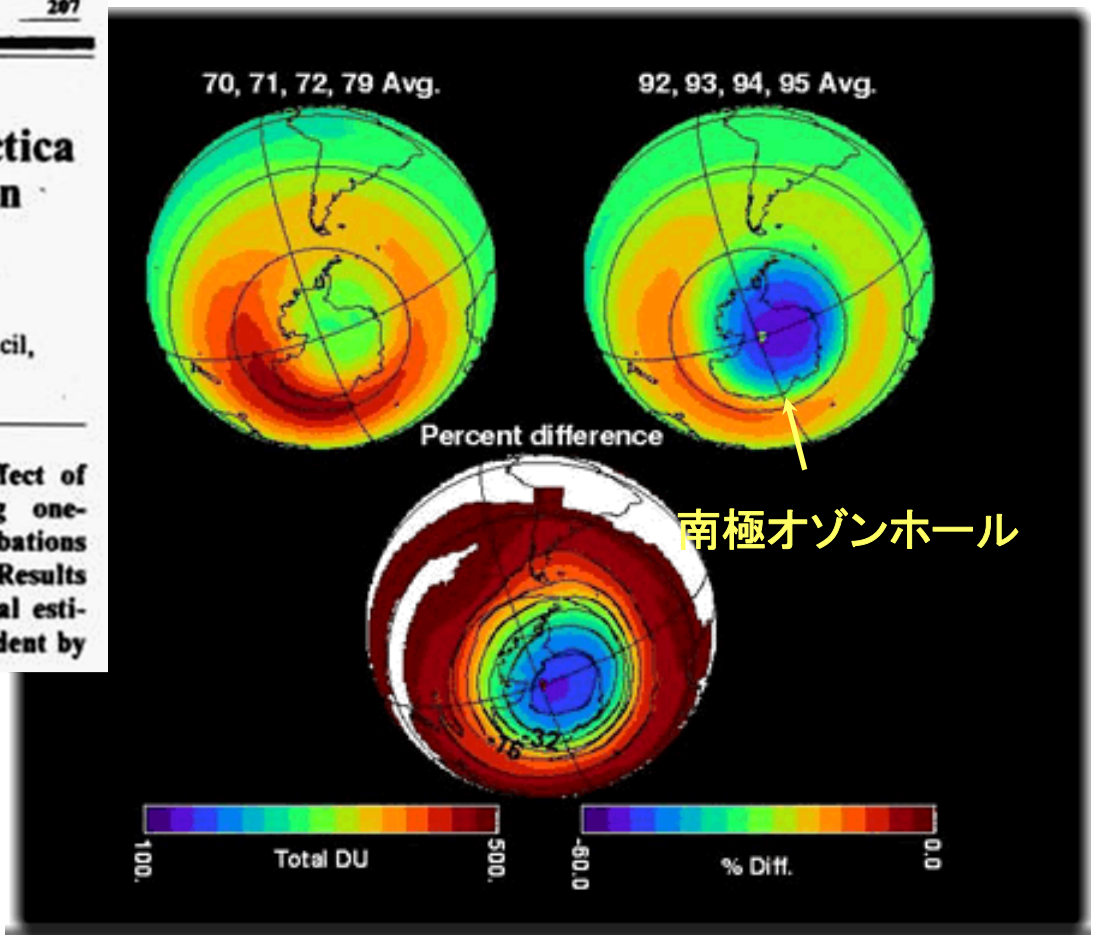
## Large losses of total ozone in Antarctica reveal seasonal ClO<sub>x</sub>/NO<sub>x</sub> interaction

J. C. Farman, B. G. Gardiner & J. D. Shanklin

British Antarctic Survey, Natural Environment Research Council,  
High Cross, Madingley Road, Cambridge CB3 0ET, UK

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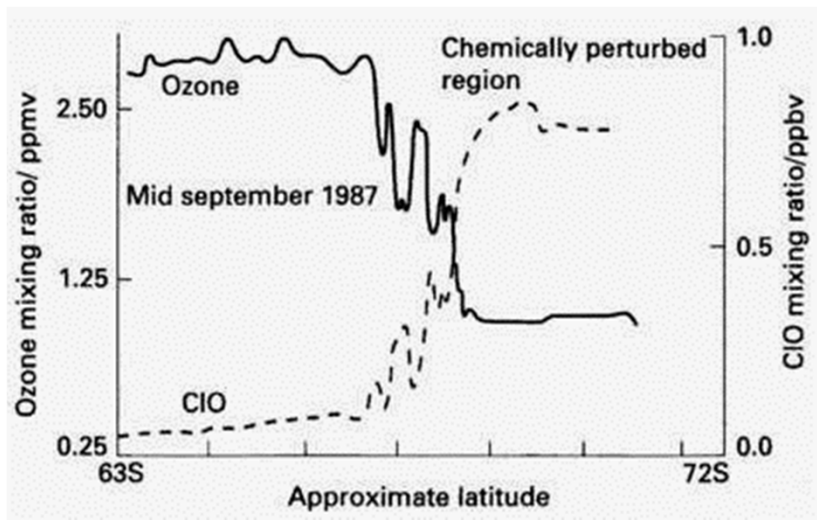
Recent attempts<sup>1,2</sup> to consolidate assessments of the effect of human activities on stratospheric ozone (O<sub>3</sub>) using one-dimensional models for 30° N have suggested that perturbations of total O<sub>3</sub> will remain small for at least the next decade. Results from such models are often accepted by default as global estimates<sup>3</sup>. The inadequacy of this approach is here made evident by



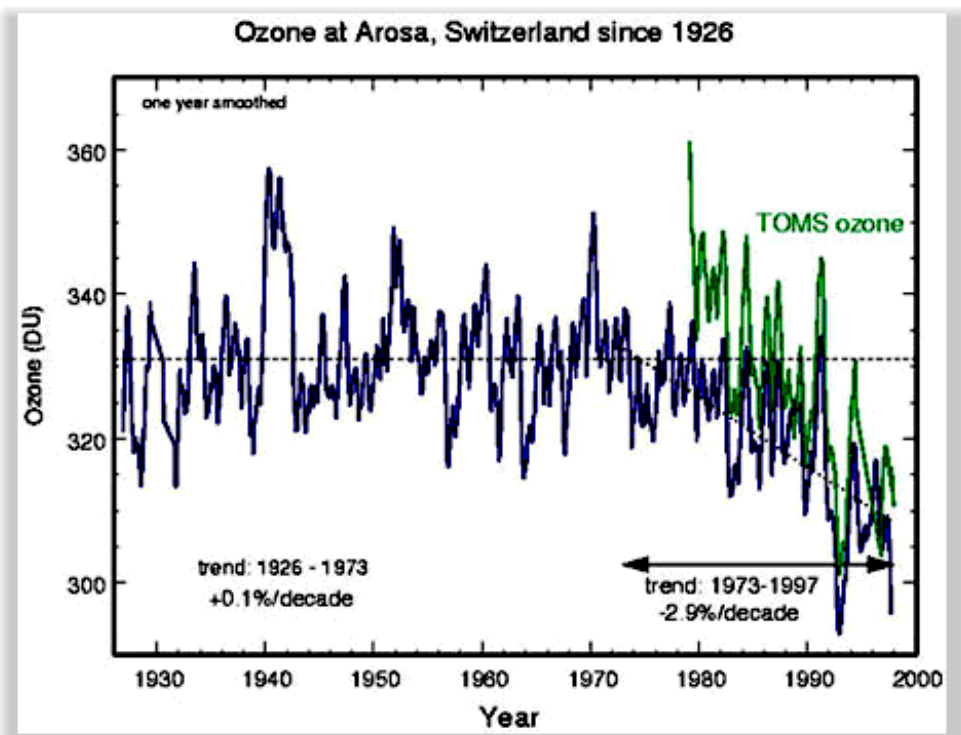
# 南極オゾンホールと CFC

1985 オゾン層保護条約  
1987 モントリオール議定書

## 中緯度成層圏オゾンの減少



南極オゾンホール端での  
ClO, O<sub>3</sub> 濃度



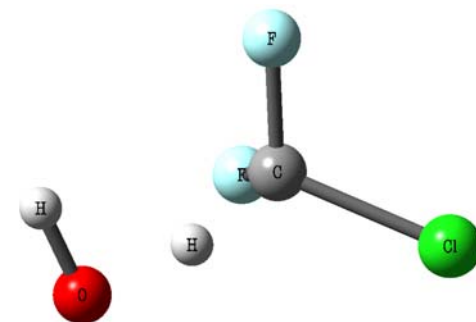
"Reaction Kinetics," M. J. Pilling and P. W. Seakins, Oxford Univ. Press, Oxford (1995).

# CFC 代替技術

## DuPont 社 - CFC代替品開発

量子化学計算によるスクリーニング

Gaussian (J. A. Pople 1998 ノーベル化学賞)



cf.) HCFC-22 と OH ラジカルの反応の遷移状態

## 対流圏(大気中)寿命

CFC, 代替品 (HCFC/HFC)	$k_{\text{OH}}$ / $\text{cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$	$\tau_{\text{OH}}$ / yr	ODP	GWP (100yr)
冷媒用				
CFC-12 ( $\text{CCl}_2\text{F}_2$ )	—	<b>102</b>	<b>1</b>	<b>8500</b>
HCFC-22 ( $\text{CHClF}_2$ )	$4.6 \times 10^{-15}$	<b>13.3</b>	<b>0.055</b>	<b>1700</b>
HFC-134a ( $\text{CH}_2\text{FCF}_3$ )	$4.2 \times 10^{-15}$	15	0	1300
発泡用				
CFC-11 ( $\text{CCl}_3\text{F}$ )	—	<b>50</b>	<b>1</b>	<b>4000</b>
HCFC-141b ( $\text{CH}_3\text{CCl}_2\text{F}$ )	$5.9 \times 10^{-15}$	<b>9.4</b>	<b>0.11</b>	<b>630</b>
HCFC-142b ( $\text{CH}_3\text{CClF}_2$ )	$3.0 \times 10^{-15}$	19.5	0.065	2000

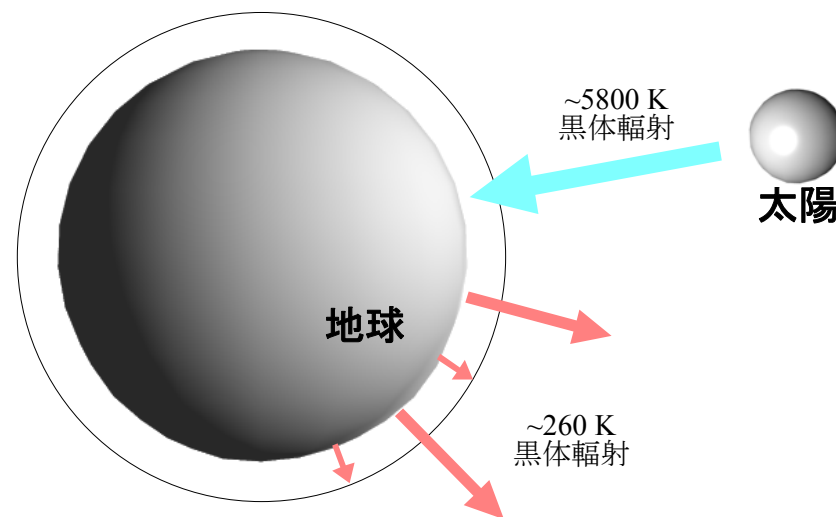
$$\tau_{\text{OH}} = 1 / ([\text{OH}]_{\text{ss}} \times k_{\text{OH}})$$

$$[\text{OH}]_{\text{ss}} \sim 5 \times 10^5 \text{ molecules cm}^{-3}$$

# 気候変動

## Climate Change

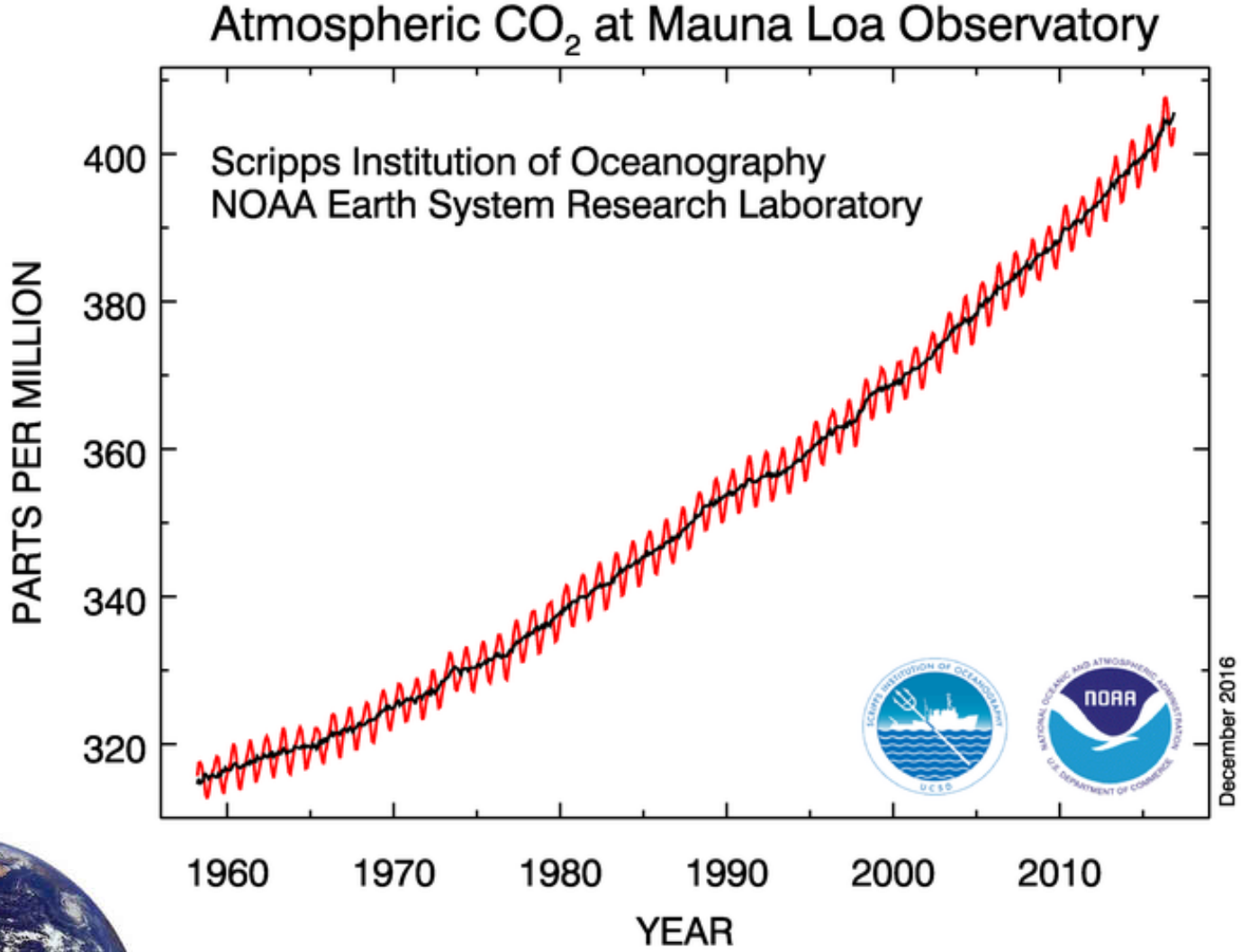
地球 = 巨大システム



# 継続観測による事実

事実としての CO<sub>2</sub> 濃度  
(~60年前から)

1958- C. D. Keeling & co-workers  
Mauna Loa Observatory (Hawaii) - 大気観測

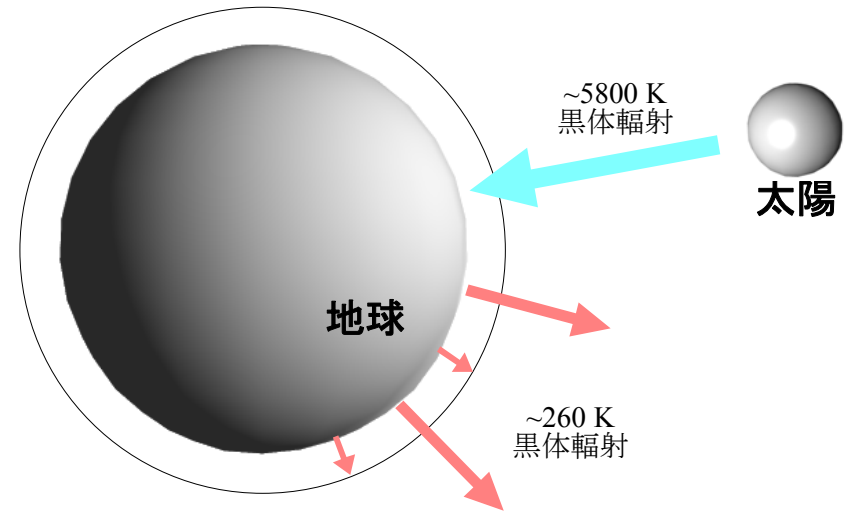


# 温室効果

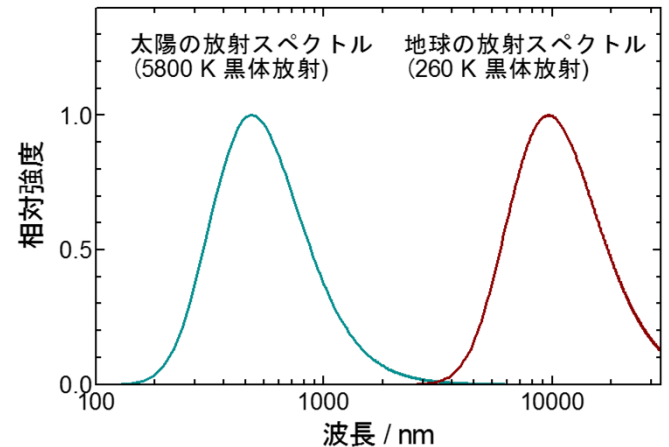
19世紀 Svante Arrhenius – CO<sub>2</sub> による温室効果  
 – “On the Influence of Carbonic Acid in the Air upon the Temperature of the Ground”  
 [Philosophical Magazine 41, 237 (1896)]

A great deal has been written on the influence of the absorption of the atmosphere upon the climate. Tyndall [2] in particular has pointed out the enormous importance of this question. To him it was chiefly the diurnal and annual variation of the temperature that were lessened by this circumstance. Another side of the question, that has long attracted the attention of physicists, is this: Is the mean temperature of the ground in any way influenced by the presence of heat-absorbing gases in the atmosphere? Fourier [3] maintained that the atmosphere acts like the glass of a hot-house, because it lets through the light rays of the sun but retains the dark rays from the ground. This idea was elaborated by Pouillet [4]; and Langley was by some of his researches led to the view, that "the temperature of the earth under direct sunshine, even though our atmosphere were present as now, would probably fall to -200 ° C., if that atmosphere did not possess the quality of selective absorption" [5]. This view, which was founded on too wide a use of Newton's law of cooling, must be abandoned, as Langley himself in a later memoir showed that the full moon, which certainly does not possess any sensible heat-absorbing atmosphere, has a "mean effective temperature" of about 45 ° C. [6]

原理: 19世紀  
 (~120年前)



収支に関わる波長の違い



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# 放射強制力への様々な寄与

Radiative forcing of climate between 1750 and 2011

Forcing agent

Confidence Level

Anthropogenic

Natural

Forcing Agent	Component	Confidence Level
Anthropogenic	Well Mixed Greenhouse Gases	Very High
	CO <sub>2</sub>	Very High
	Other WMGHG	Very High
	CH <sub>4</sub>	Very High
	N <sub>2</sub> O	Very High
	Halocarbons	Very High
	Ozone	High
	Stratospheric	High
	Tropospheric	High
	Stratospheric water vapour from CH <sub>4</sub>	Medium
Anthropogenic	Surface Albedo	High/Low
	Land Use	High/Low
	Black carbon on snow	High/Low
	Contrails	Medium
	Contrail induced cirrus	Low
Anthropogenic	Aerosol-Radiation Interac.	High
	Aerosol-Cloud Interac.	Medium
	Aerosol-Cloud Interac.	Low
Total anthropogenic	Total anthropogenic	Medium
	Solar irradiance	Medium

放射強制力への多様な寄与

増加を続ける CO<sub>2</sub>(二酸化炭素), CH<sub>4</sub>(メタン), N<sub>2</sub>O(亜酸化窒素)

Top Graph: CO<sub>2</sub> (ppm) and pH<sub>total</sub>. Legend: CO<sub>2</sub> - MLO (ppm), CO<sub>2</sub> - SPO (ppm), pCO<sub>2</sub> - HOT (μatm), pH<sub>total</sub> - HOT.

Middle Graph: ΔO<sub>2</sub> (ppm) and CH<sub>4</sub> (ppb). Legend: ΔO<sub>2</sub> - ALT, ΔO<sub>2</sub> - CGO, CH<sub>4</sub> - MLO, CH<sub>4</sub> - SPO.

Bottom Graph: δ<sup>13</sup>C (permil) and N<sub>2</sub>O (ppb). Legend: δ<sup>13</sup>C - MLO, δ<sup>13</sup>C - SPO, N<sub>2</sub>O - MHD, N<sub>2</sub>O - CGO.

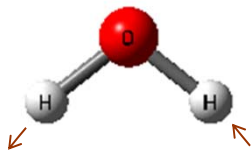
Climate Change 2013 - The Physical Science Basis  
 Working Group I Contribution to the Fifth Assessment Report of the IPCC  
 available at <http://www.ipcc.ch/>



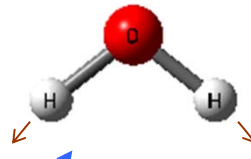
# 温室効果気体

## 大気の赤外吸収スペクトル (光路長 10 cm)

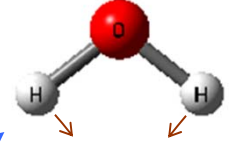
分子の振動と回転に起因する複雑なスペクトル



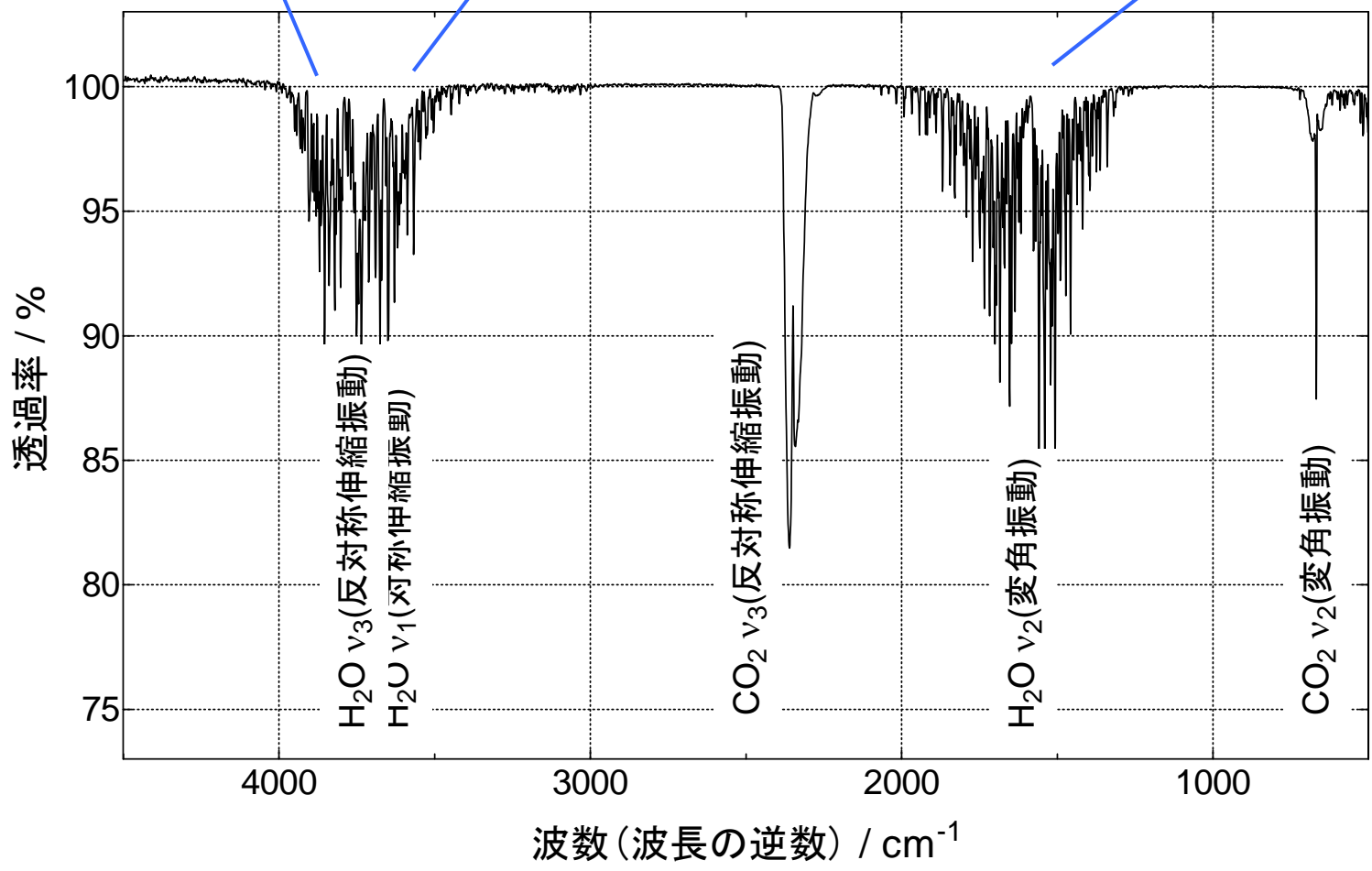
反対称伸縮振動



対称伸縮振動



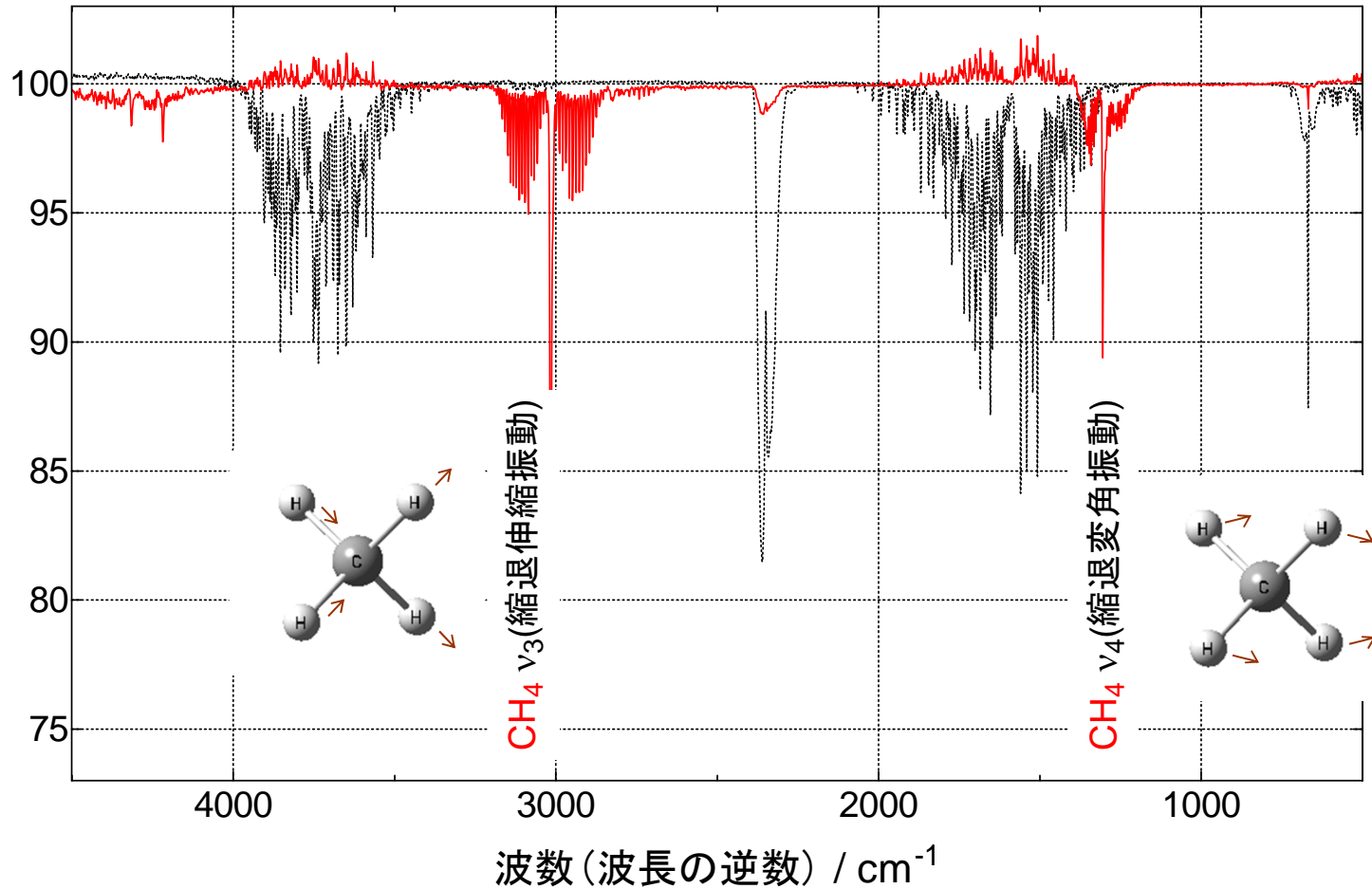
変角振動



# メタン - 温室効果気体

H<sub>2</sub>O, CO<sub>2</sub>の吸収のない  
大気の「窓」に吸収をもつ → 温室効果大

## メタン・大気の赤外吸収スペクトル



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# オゾン - 温室効果気体

成層圏のオゾン ... ○ 紫外線を遮蔽  
 対流圏のオゾン ... × 温室効果気体

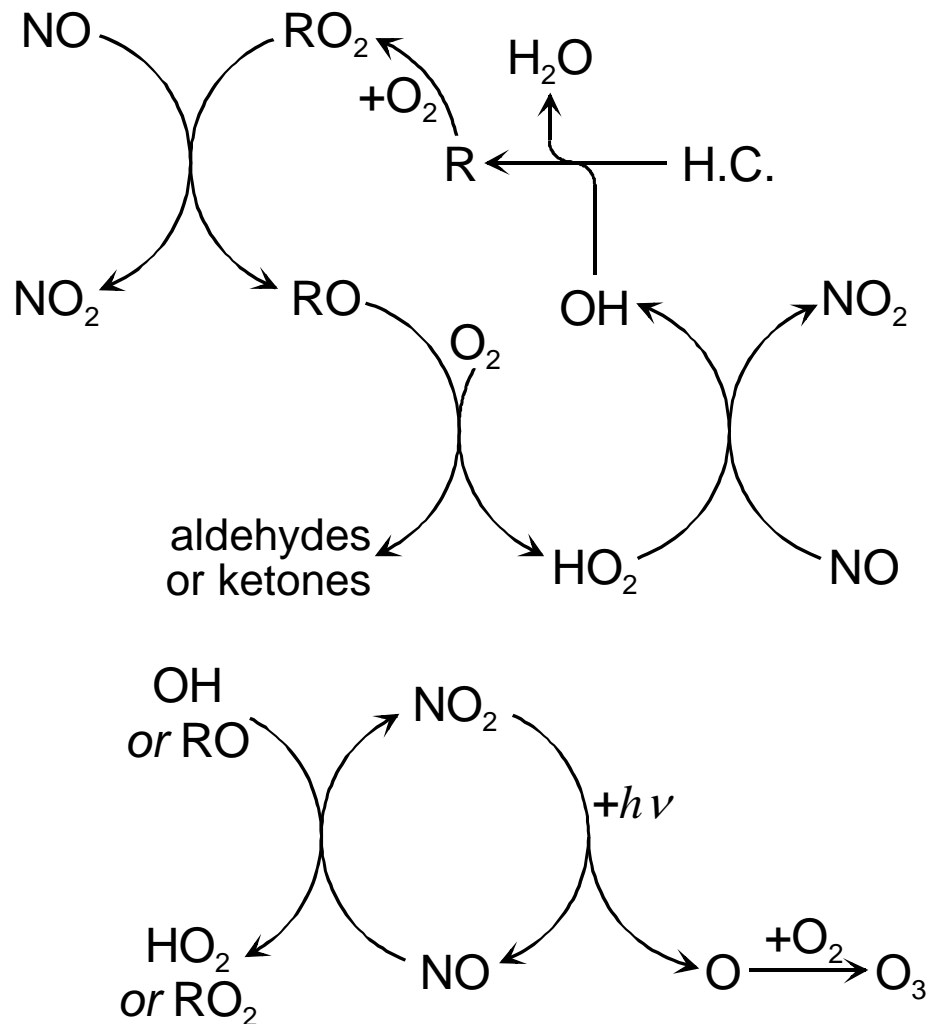
## オゾン・大気の赤外吸収スペクトル

The figure displays the infrared absorption spectrum of ozone ( $O_3$ ). The vertical axis represents transmittance, ranging from 75 to 100. The horizontal axis represents wavenumber in  $cm^{-1}$ , ranging from 4000 to 1000. The spectrum shows several characteristic absorption bands:

- $O_3 \nu_1 + \nu_3$  (around 2100  $cm^{-1}$ )
- $O_3 \nu_1$  (対称伸縮振動) (around 1100  $cm^{-1}$ )
- $O_3 \nu_3$  (反対称伸縮振動) (around 1100  $cm^{-1}$ )
- $O_3 \nu_2$  (変角振動) (around 700  $cm^{-1}$ )

The spectrum is overlaid with a blue line that highlights the absorption bands. The background is a grid of dotted lines.

# 対流圏のオゾン



HC(炭化水素) / NO<sub>x</sub> ← 大気汚染

**光化学スモッグ**

**O<sub>3</sub> (オキシダントの主成分)**

HC 酸化 ... XO<sub>2</sub> (HO<sub>2</sub>/RO<sub>2</sub>)



**OH ラジカル**

**(大気の掃除屋/洗浄剤)**

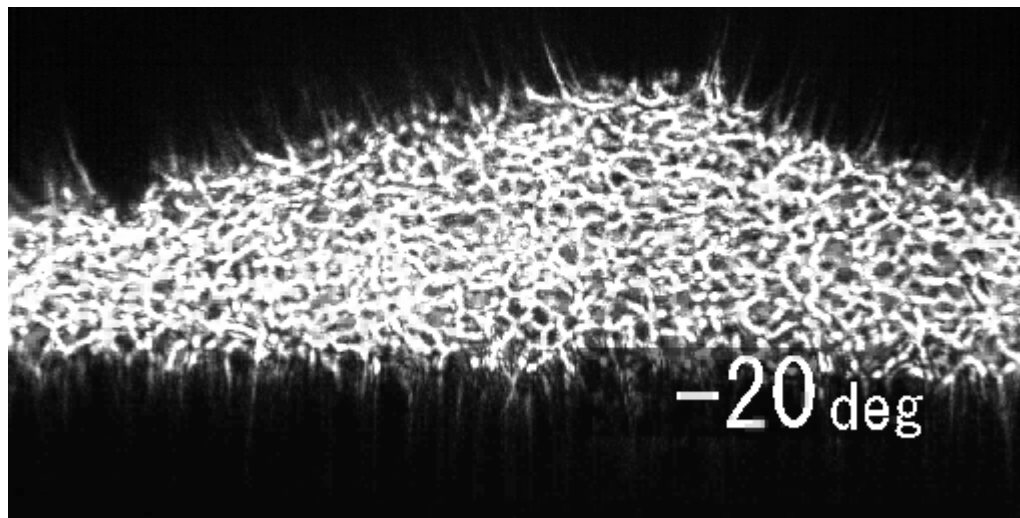
**の発生源**

(detergent of the atmosphere)

自然起源 NO<sub>x</sub> ← 土壌/雷

# 燃焼技術

## combustion technology



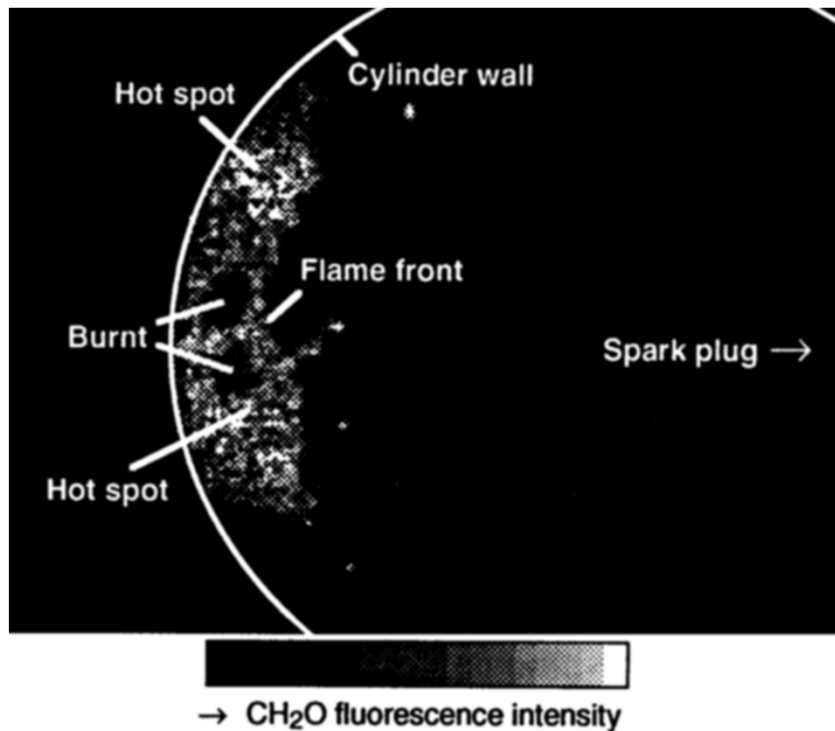
ガソリンエンジン  
内の可視化映像

by courtesy of  
Dr. M. Kaneko

# エンジンノック — 課題

## — 現象

- 自動車エンジンで発生する  
好ましくない 自着火



J. Warnatz, U. Maas, and R. W. Dibble,  
"Combustion," Springer, Berlin, 1996.

## — 問題

- 機関の致命的 損傷
- 抑制 (低圧縮比化/着火時期遅延)  
は 熱効率の低下 を伴う



This photo of a badly damaged piston indicates the effects of long-term engine knock.

Lawrence Livermore National Laboratory

Science and Technology Review, Dec. 1999, Lawrence Livermore National Laboratory, UCRL-52000-99012 (1999).

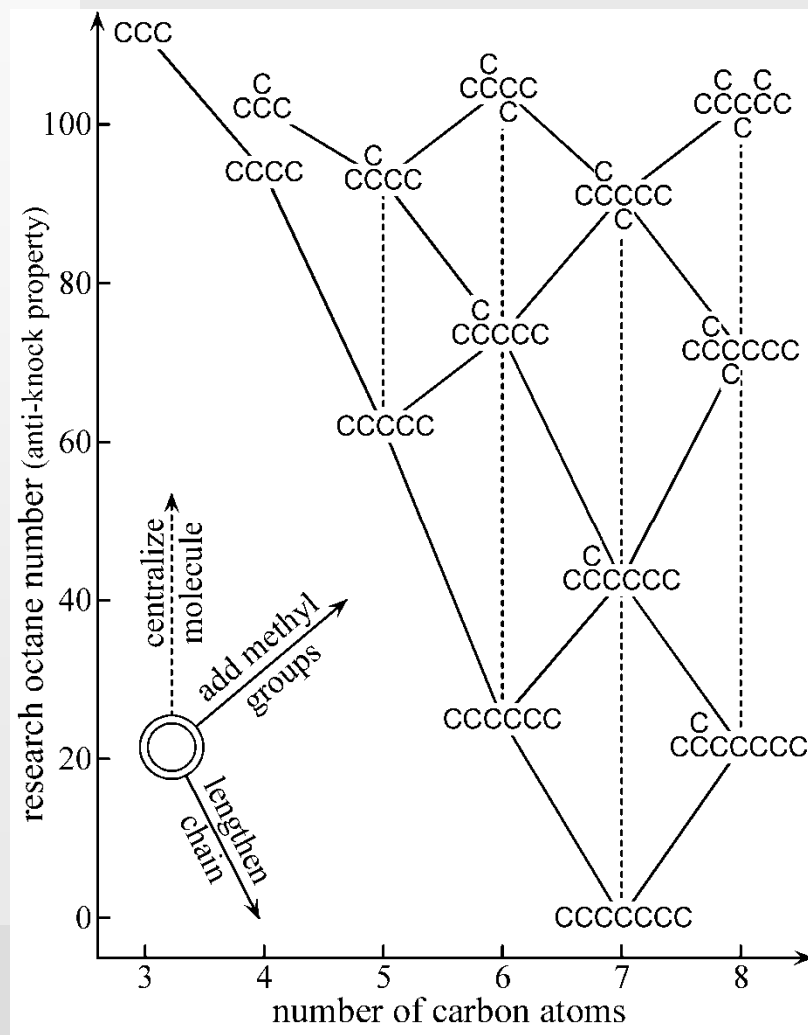
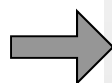
# オクタン価

## ー ガソリンの 耐ノック性の指標

- オクタン価が高い  
= 自着火しにくい  
(良質なガソリン)
- 燃料炭化水素の化学構造  
と明瞭な関係がある

→ どうしてか？

→ どう 制御するのか？



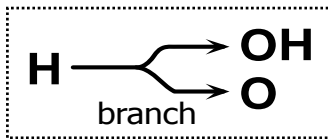
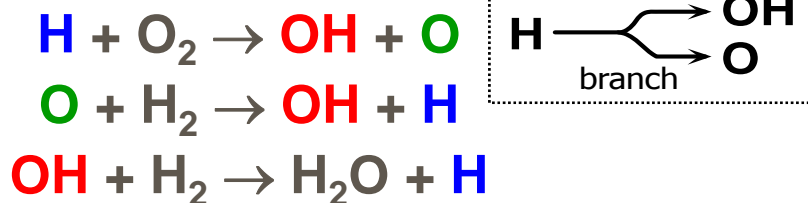
based on: W. G. Lovell, *Ind. Eng. Chem.*, 40, 2388 (1948).





# 連鎖反応 — 分岐連鎖反応

## H<sub>2</sub>-O<sub>2</sub> 一分岐連鎖反応 (連鎖着火)



活性種は消えない (右辺に残る)

- 連鎖担体 (H, O, OH) 自己増殖  
→ 自己加速 → 連鎖着火

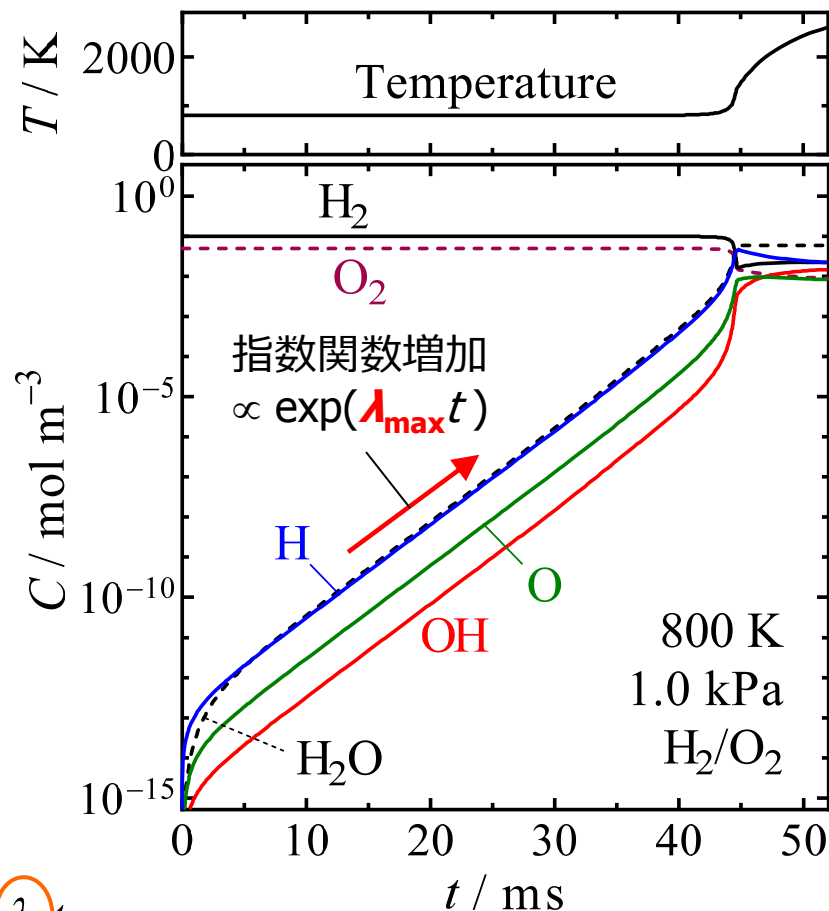
$$\begin{aligned} x &= [\text{H}], y = [\text{O}], z = [\text{OH}], \\ R_1 &= k_1[\text{O}_2], R_2 = k_2[\text{H}_2], R_3 = k_3[\text{H}_2] \end{aligned}$$

$$\begin{pmatrix} x' \\ y' \\ z' \end{pmatrix} = \begin{pmatrix} -R_1 & R_2 & R_3 \\ R_1 & -R_2 & 0 \\ R_1 & R_2 & -R_3 \end{pmatrix} \begin{pmatrix} x \\ y \\ z \end{pmatrix}$$

$$\mathbf{x} = \sum_i a_i \mathbf{s}_i e^{\lambda_i t}$$

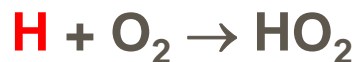
ヤコビ行列

温度/圧力/組成 が一定 (着火誘導期中) → 不変



$\lambda_{\max} > 0$  ... 発散項

# 自着火限界 $\leftrightarrow$ 最大固有値 = 0



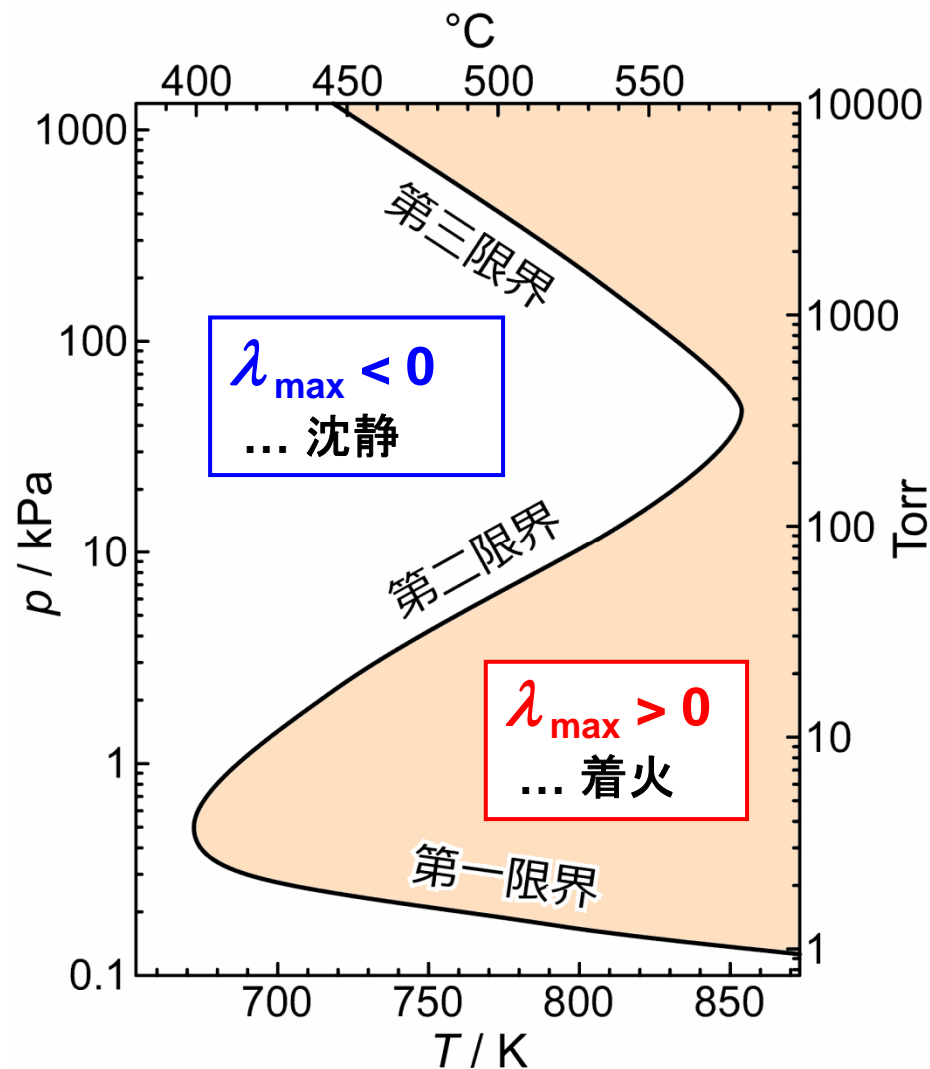
$$\begin{pmatrix} x' \\ y' \\ z' \end{pmatrix} = \begin{pmatrix} -R_1 - R_4 & R_2 & R_3 \\ R_1 & -R_2 & 0 \\ R_1 & R_2 & -R_3 \end{pmatrix} \begin{pmatrix} x \\ y \\ z \end{pmatrix}$$

ヤコビ行列

$$\mathbf{x} = \sum_i a_i \mathbf{s}_i e^{\lambda_i t}$$

第二限界はこのヤコビ行列で説明可能

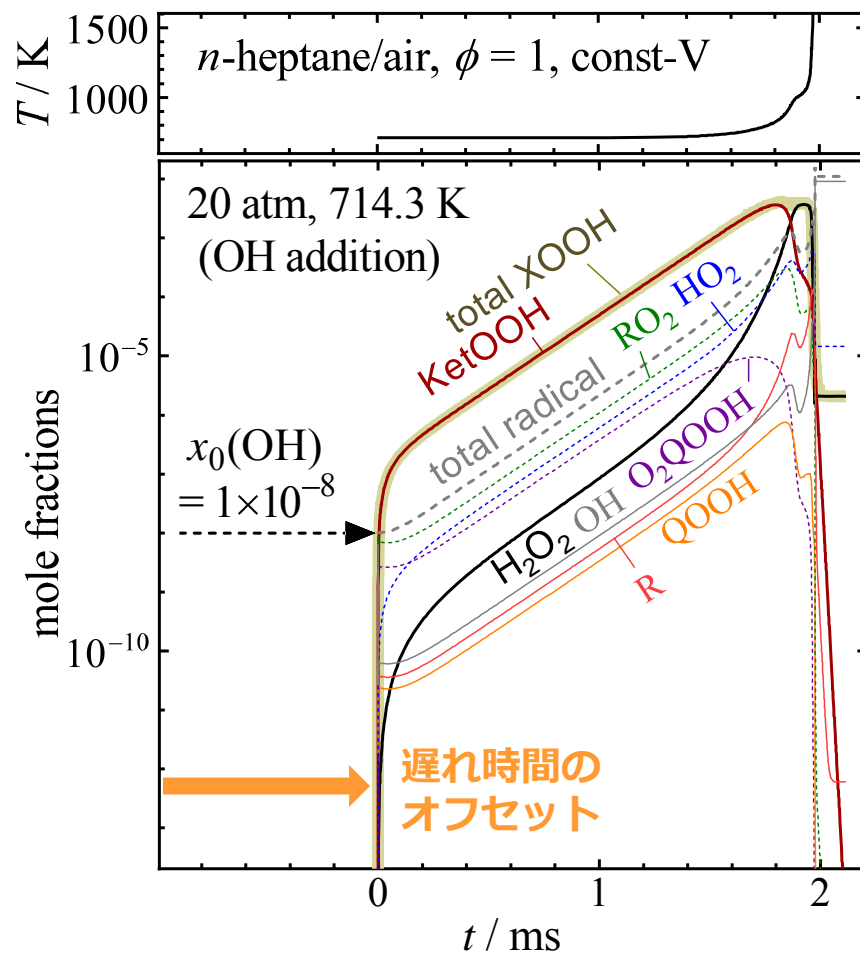
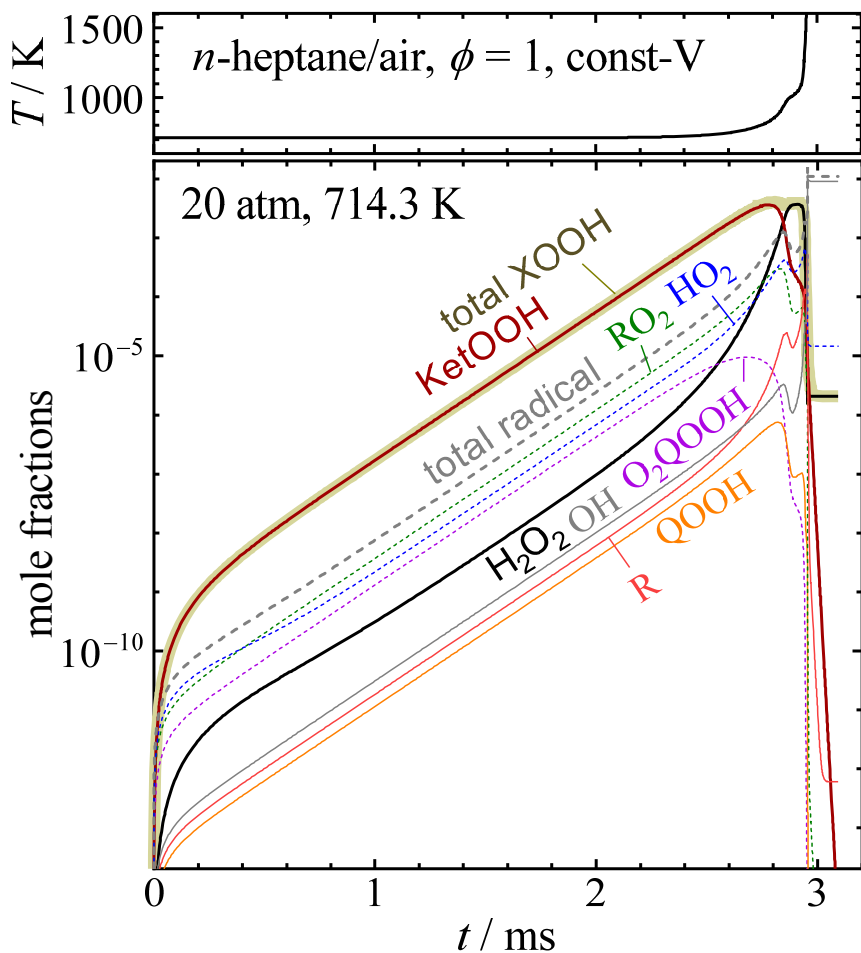
$\text{H}_2:\text{O}_2 = 2:1$  混合気の自着火限界



# 活性種添加と定常状態

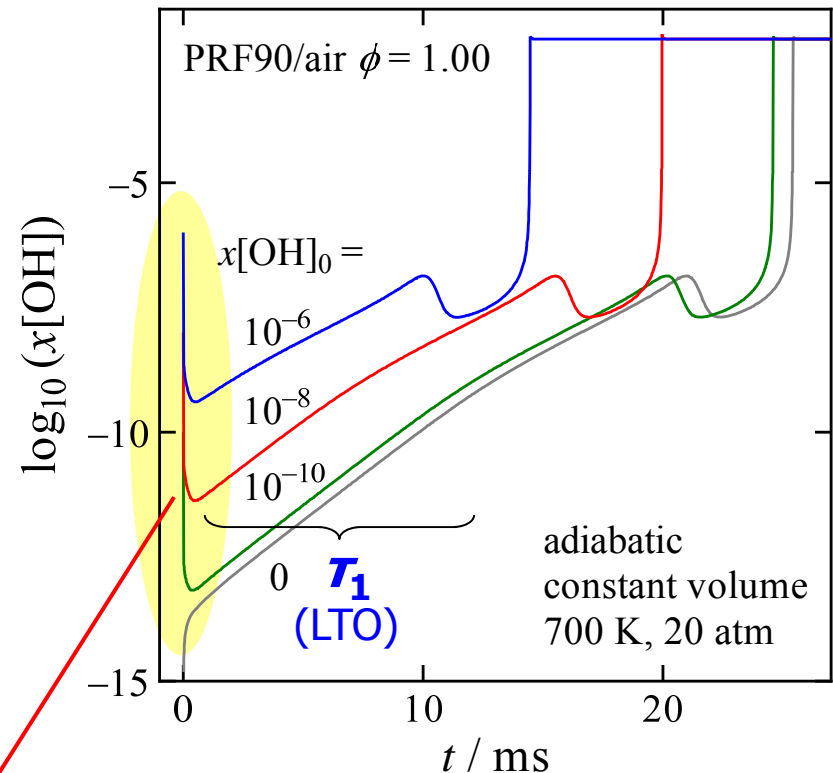
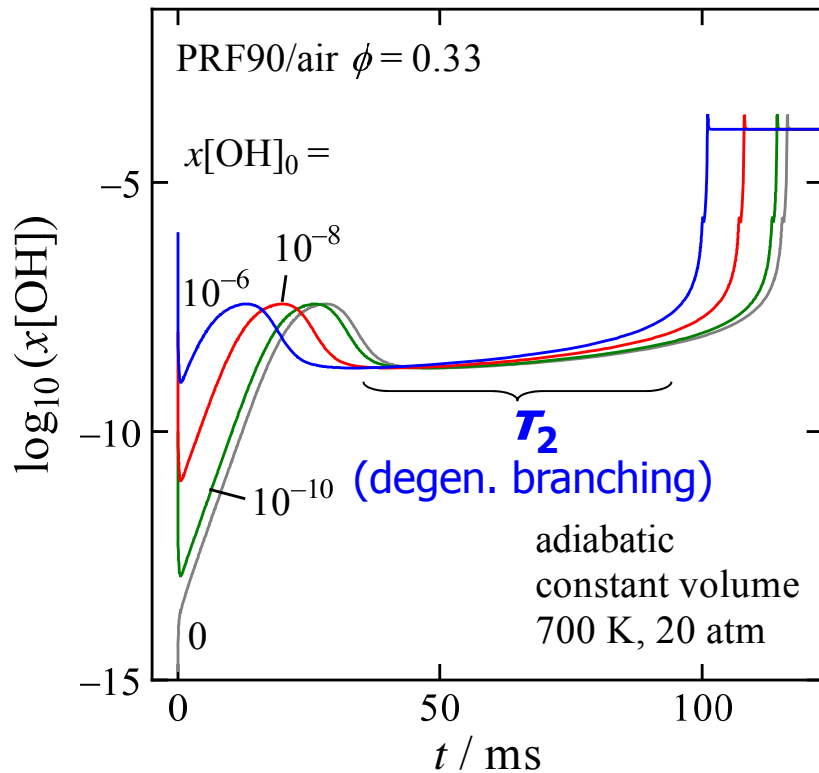
- 活性種添加量  $\approx$  全活性種増加量
- OHを添加しても速やかに定常濃度になる
- 添加効果は時間に指数関数的に減少

- 活性種によらない
- 頑健な定常状態



# 放電の「対数的」効果

- ラジカルの添加は指数関数的増加の始点を変える → 効果は対数的
- $\tau_1$  (冷炎着火遅れ) のみに影響する
- $\tau_2$  が支配的な場合には効果が小さい (例えば低当量比)



$$[\text{OH}]_{\text{ss}} \approx 10^{-3} [\text{OH}]_{\text{added}}$$

# まとめ

## — オゾン層破壊

- バックグラウンド(非汚染)大気中の CFC (~1971)
- 連鎖反応の可能性の示唆 (1974)
- 南極環オゾンホール (1985)
- → 規制・代替技術 ... 「優等生」

## — 気候変動

- 温室効果の原理 (19世紀)
- バックグラウンド CO<sub>2</sub> の増加 (1958~)
- 気候変動の証明は困難 ... 「劣等生？」

## — 燃焼技術

- 燃焼技術 = エネルギー技術
- 数千~ 数万~ の素過程を用いたモデリング = システムの科学
- → 制御の可能性探索